

1           **Supporting Information for "Observations of near-surface**  
2           **current shear help describe oceanic oil and plastic transport"**

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## 21 Introduction

22 The text that follows describes each type of sensor used in acquiring the data shown  
23 in the main manuscript. The figures provide additional information which may be useful  
24 to the reader.

### 25 Text S1 - Polarimetric Camera

26 A bow tower-mounted polarimetric camera was used to measure the spatiotempo-  
27 ral development of the short wave slope field [Zappa *et al.*, 2008], taking into account the  
28 vessel's rotational [Laxague *et al.*, 2015] and translational [Laxague *et al.*, 2017] motion.  
29 Each ten seconds of sea surface slope fields were packaged into a 780x780x200 array and  
30 subjected to a 3D Fourier transform, producing a wavenumber-frequency spectrum; these  
31 were averaged over the sampling duration of four hours. Wavenumbers ranged from 5-  
32 1500 rad/m; frequencies ranged from 0.1-10 Hz. Bound waves [Plant *et al.*, 1999] were  
33 identified and removed from the spectra. Once this operation had been performed, depar-  
34 tures from the gravity-capillary linear dispersion relation were logged at each wavenum-  
35 ber component and used to estimate current [Laxague *et al.*, 2017], with depth taken from  
36 literature [Plant and Wright, 1980] as 4.4% of the wavelength. All measured velocity  
37 profiles were arranged to yield a profile (and the 5<sup>th</sup>-95<sup>th</sup>-percentile range) that was bin-  
38 averaged in the vertical to produce layer thicknesses of 1 mm (as shown in main document  
39 Figure 2).

### 40 Text S2 - Drifting Instruments

41 The GPS-tracked drifters ("CARTHE" drifters- [http://www.greenwaveinstruments.](http://www.greenwaveinstruments.com)  
42 [com](http://www.greenwaveinstruments.com)) were specifically developed with goals of being lower cost, easier to handle, and  
43 more biodegradable than previous designs- all while still accurately following the near-  
44 surface Lagrangian velocity. Each drifter body was constructed from injection-molded  
45 bio-plastic (PHA), biodegradable in seawater, and low-toxicity electronics for data trans-  
46 mission. A single drifter combines a surface float (containing the GPS transmitter and  
47 batteries) and a submerged drogue (draft of 70 cm, center of drag at 30 cm) connected  
48 by a chain tether that flexes to allow the drifter to more accurately follow the slope of the  
49 waves at the surface. Its drifting characteristics were extensively quantified in a wind-wave  
50 tank and in the field [Novelli *et al.*, 2017]. Measurements show that the drifter (respec-

51 tively the floater alone) follows the average Lagrangian current, including wave Stokes  
 52 drift, in the upper 0.65 m (respectively 0.035 m), to an accuracy better than 0.5% (re-  
 53 spectively 2%) of the equivalent 10 m neutral wind speed  $U_{10}$ , over a range of 8-23 m/s.  
 54 In field trials with  $U_{10}$  of 5-15 m/s, the velocities of CARTHE drifters matched that of  
 55 neighboring CODE drifters to within 0.1% of  $U_{10}$ . Biodegradable bamboo dinner plates  
 56 were deployed via small boat and tracked to measure near-surface currents. Each plate had  
 57 a draft of 0.015 m and a diameter of 0.28 m, floating just below the surface. Given typ-  
 58 ical dry and wet bamboo densities of  $860 \text{ kg/m}^3$  and  $1160 \text{ kg/m}^3$ , respectively [*Amada*  
 59 *and Lakes*, 1997]- and that the nearest-surface water density during sampling time was  
 60  $\approx 1140 \text{ kg/m}^3$ , we estimate that the damp plates were at least 95% submerged, exposing  
 61 a negligible amount to direct wind forcing. Tracking was performed on images recorded  
 62 using a gimbaled high-resolution (20 Megapixels) camera attached to a commercial UAV  
 63 (drone). Each frame was georeferenced using the drone's GPS, altitude, and camera atti-  
 64 tude records. At an altitude of 130 m, each plate was represented by approximately  $10 \times 10$   
 65 pixels. In order to minimize the effect of error introduced into the current measurements  
 66 from uncertainties in the GPS record (e.g., as described in [*Miyao and Isobe*, 2016]), the  
 67 leading edge of the bamboo plates was tracked at the beginning and end of each drone  
 68 flight. This ensured that the  $\approx 5$  meter total error in drone position between the two frames  
 69 was diluted by a horizontal scale of  $\approx 500$  meters, or  $\approx 1\%$  error. This single drift speed  
 70 was found to be within 2-3% of the drift speed computed from consecutive UAV frames,  
 71 indicating that GPS error from one drone position to the next did not appreciably contami-  
 72 nate the current estimates.

### 73 **Text S3 - ADCPs**

74 Currents in the upper 5m of the water column were measured by a REMUS-100  
 75 Autonomous Underwater Vehicle (AUV) deployed concurrently with the drifter and plate  
 76 release experiment. For the most of its mission, the AUV navigated at 2m depth along the  
 77 parallel lines underneath the drifter field, with occasional dives to 10m to sample vertical  
 78 stratification. The vehicle was equipped with upward- and downward-looking 1200kHz  
 79 RDI ADCP. For the experiment, the ADCP was operated in the pulse-coherent mode, with  
 80 the vertical bin resolution of 7 cm over the range of  $\pm 2.1$  m above and below the vehicle  
 81 (except for the  $\pm 0.4$ m blanking distance gap in the middle). Velocity measurements were  
 82 referenced to the bottom, tracked continuously using dedicated ADCP pings. ADCP sam-

83 pling rate was on the order of 0.65 Hz. Estimated RMS single-ping accuracy of velocity  
 84 measurements was 1.8 cm/s. 80-minute average current profile spanning depth range from  
 85 0.2 to 4.4 m was used in the study (Fig. 2), with an estimated standard error of <2 cm/s  
 86 attributable to wave orbital motions with the amplitude of up to 20 cm/s near the surface.  
 87 Currents below 4 m were measured with the pole-mounted 1200kHz RDI ADCP on the  
 88 R/V F. G. Walton Smith. The instrument was positioned in a downward-looking orienta-  
 89 tion at 3.25 m depth with a 50 cm blanking distance. Measurements spanned the range  
 90 from 4 m to 1 m off the bottom, with 50 cm vertical bin spacing. Velocities were pro-  
 91 cessed using the UHDAS system [Firing *et al.*, 2012], referenced to the bottom (correcting  
 92 for ship translation) and time-averaged over the duration of the experiment.

#### 93 **Text S4 - Wind, Waves, Stokes Drift, and Water Density**

94 The wind velocity vector was recorded at 20 Hz from a tower-mounted Campbell  
 95 Scientific IRGASON sonic anemometer. Water surface elevation was recorded at 20 Hz  
 96 from a bow-mounted pentagonal array of acoustic altimeters. With the aid of a six degree-  
 97 of-freedom inertial motion unit, both sets of data were corrected for ship motion [Ancil  
 98 *et al.*, 1994]. The wind stress vector (and through it, the friction velocity  $u_*$ ) was com-  
 99 puted from the corrected wind velocity time series via the eddy covariance method [Edson  
 100 *et al.*, 2013]. The earth-referenced array of water surface elevation measurements were  
 101 used to calculate the gravity wavenumber directional spectrum [Donelan *et al.*, 1996].  
 102 This was combined with the high-wavenumber spectrum computed from the polarimetric  
 103 slope fields described above and converted into an intrinsic frequency directional spectrum  
 104 (making sure to conserve energy [Plant, 2009] and further utilized to compute the direc-  
 105 tional Stokes Drift profile. The particular technique used for this computation ("2Dh-SD")  
 106 accounts for the directional spreading of the wave field [Webb and Fox-Kemper, 2015].  
 107 The density profile shown in Figure 1c was computed using the International Equation of  
 108 State for sea water (EOS-80) via temperature and salinity measurements obtained from a  
 109 series of Castaway CTD profiles performed throughout the experiment.

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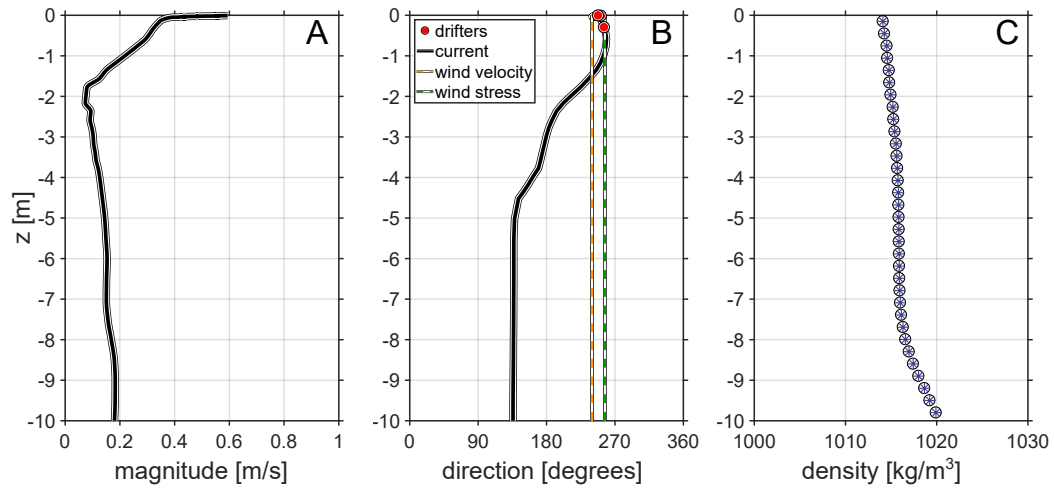
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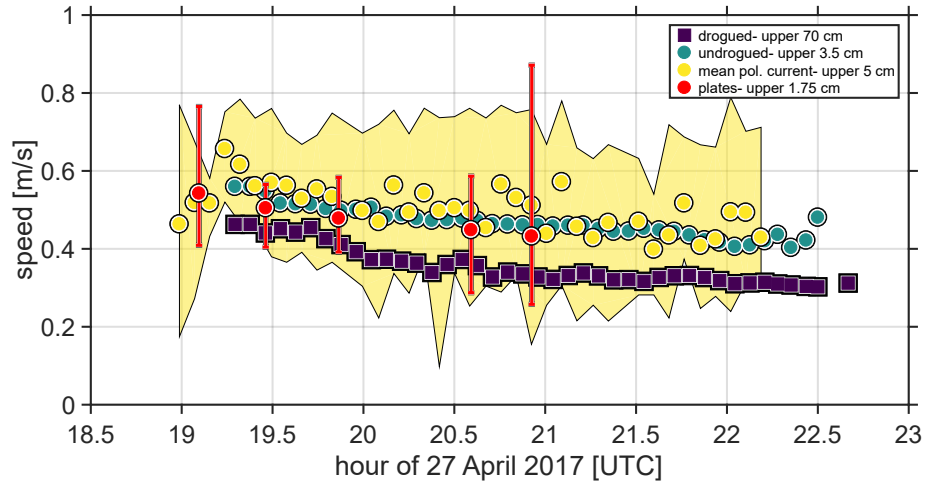
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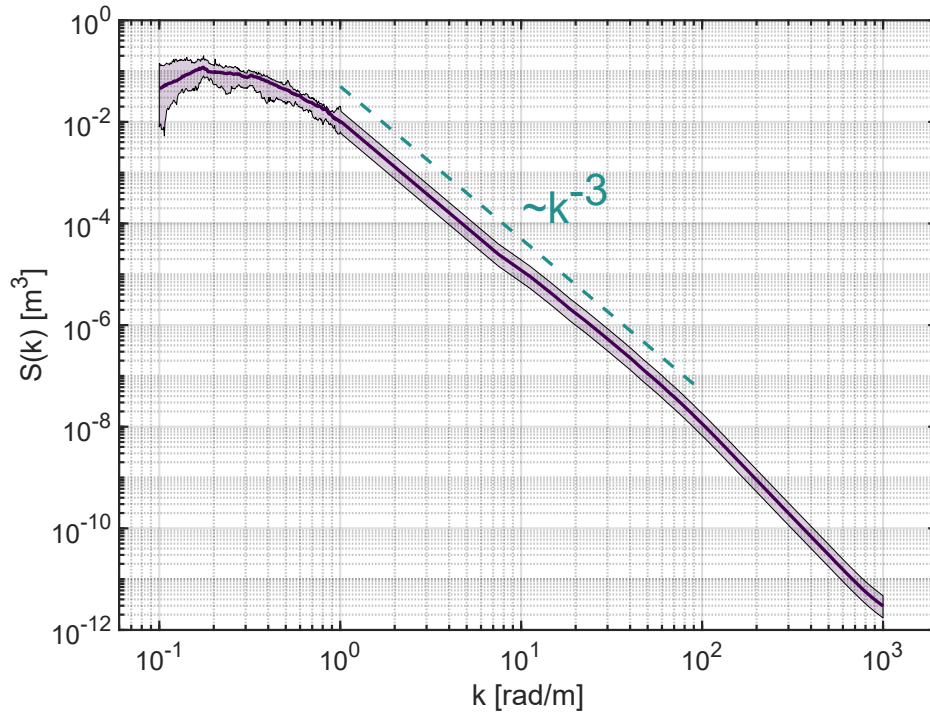


156 **Figure 1.** Current and density profile: Magnitude (a) and direction (b) of observed current, bearing of  
 157 drifting instruments, and orientation of wind velocity & stress vectors. (c) Profile of water density obtained  
 158 from the Castaway CTD profiler.

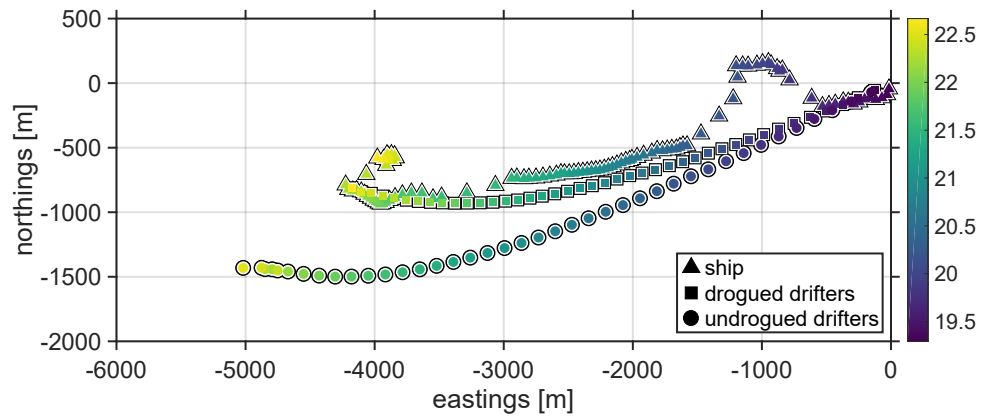




159 **Figure 2.** Current time series: Time series of drifter/plate centroid translation speeds and mean polari-  
 160 metric camera-determined current magnitude. Yellow shaded region (red bars) represents 5<sup>th</sup>-95<sup>th</sup> percentile  
 161 range for the camera-derived current magnitude (plate speed).



162 **Figure 3.** Composite spectrum computed from ocean surface gravity waves as measured by the bow-  
163 mounted acoustic altimeter array and the polarimetric camera, defined from long gravity waves to pure  
164 capillary waves. Used for estimation of Stokes drift profile. Shaded region represents the 5<sup>th</sup>-95<sup>th</sup> percentile  
165 range. Dashed line shows the  $k^{-3}$  slope in log-log space.



166 **Figure 4.** Trajectories of ship (triangles), drogued drifter centroid (squares), and undrogued drifter cen-  
167 troid (circles) during observation time. Origin of local coordinate system is deployment location. Colorbar  
168 indicates the hour (UTC) of April 27th, 2017.